



重新审视深层油气成藏模式： 以塔里木盆地为例

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内容提要:近年来,随着油气勘探不断向深层—超深层领域拓展,深层展现出了巨大的油气勘探潜力,同时也在成烃—成储—成藏等方面浮现出一系列科学问题。本文在广泛调研国内外相关研究的基础上,梳理了深层油气成藏环境和生、储、盖成藏要素的特殊性,重点讨论了深层油气藏在形成与演化过程中需要深入关注的四个基本问题:① 深层中的油气生成、储集空间形成、油气相态和运移等研究,均需要以物质守恒和能量守恒基本定律为前提开展;② 深层经历了盆地演化全过程,需要从动态演化角度研究油气成藏;③ 需要探索新的实验方法,加大对深层液态烃稳定性与保存深度下限的研究;④ 注重多学科融合与多技术交叉,解决深层复杂的地质问题。在此基础上,提出了深层油气藏最为可能的两种成藏模式:① 中—浅层油成藏、深埋保持型;② 长期浅埋、晚期快速深埋(凝析)气成藏型。以塔里木盆地台盆区顺北地区和库车坳陷博孜—大北地区为研究对象,应用储层地球化学分析、流体包裹体系列分析技术、方解石原位 U-Pb 定年技术和盆地模拟技术,对两个地区油气成藏模式进行研究。结果表明:顺北地区奥陶系深层油气藏为“早期中—浅层成藏、后期持续深埋保存”的成藏模式,油气成藏后相对稳定的构造背景是油气藏能保持至今的关键因素;博孜地区白垩系现今超深层天然气藏的形成主要发生在深层至超深层,“长期浅埋、晚期快速深埋”的构造—埋藏演化背景是该类型深层油气藏形成的关键,进一步佐证了深层油气藏的两种成藏模式。研究成果在深化深层油气成藏理论认识和指导深层油气勘探方面具有一定意义。

关键词:深层油气成藏模式;成藏年代学;流体包裹体;盆地模拟;塔里木盆地

20 世纪 50~60 年代提出的干酪根热裂解生油气理论是石油地质领域具有里程碑意义的研究成果,有效指导了之后 60 余年的油气勘探。该理论认为,沉积岩中的干酪根在 50~115℃ 开始生成石油,生油带主要位于深度小于 4000 m 的中浅层,而当温度大于 150~160℃ 时,原油便开始进入裂解阶段(Tissot and Welte, 1984; Quigley and MacKenzie, 1988)。据此,传统“地球能源黄金带”理论提出,世界上 90% 的石油和天然气储藏在温度为 60~120℃ 的地层中(Buller et al., 2005; Nadeau, 2011)。20 世纪 90 年代,Macgregor(1996)对全球 350 个已发现的巨型油藏研究也发现,全球工业油藏的年龄中

值约为 35 Ma,且绝大部分油藏分布在埋深<4500 m 的中浅层。进入 21 世纪,随着油气勘探的不断深入,全球深层(埋深>4500 m)、超深层(埋深>6000 m)油气勘探进程与成果呈快速增长趋势(贾承造和庞雄奇,2015)。据 IHS 公司统计,截至 2020 年底,全球共发现埋深大于 4500 m 的油气田(藏)1975 个,大于 6000 m 的油气田(藏)285 个,证实了深层油气勘探的重要地位(匡立春等,2021)。我国陆上深层—超深层油气勘探领域主要集中在中—西部叠合盆地,据统计,陆上深层石油资源占陆上石油总资源量的 28%,陆上深层天然气资源占陆上天然气总资源量的 52%(中国石油学会,2016)。塔里木盆地是

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我国最大的陆上含油气盆地(田军等,2021),目前油气勘探在时间和空间维度上正向着更古老、更深层进军,如,近年来,台盆区轮探1井在8203~8260 m的寒武系储层获得轻质原油(杨海军等,2020)、塔北隆起南坡富源1井在7711 m的奥陶系储层获得工业油流(Zhu Guangyou et al.,2018)、库车坳陷博孜9井在7600~7880 m深度试油获高产工业油气流(田军等,2020)。据第四次资源评价,塔里木盆地超深层石油资源量为 34.5×10^8 t、天然气资源量为 5.98×10^{12} m³,分别占盆地石油和天然气资源量的46%和51%,资源探明率低,勘探潜力巨大(黄少英等,2018;杨学文等,2021)。

叠合盆地深部地层往往经历多期构造-沉积演化历史,发育多套烃源岩、多类型储层和多套生储盖组合(赵文智等,2015),含油气盆地在复杂地质演化历史中,温度场和压力场相互作用、不断变化,控制着油气生成与相态、储层的成岩与孔隙演化(邱楠生等,2018),造成油气成藏过程与模式难以准确重建。前人在对不同含油气盆地研究基础上,认为深层油气往往具有多期次成藏和多过程改造的特征(赵孟军和张水昌,2004;罗晓容等,2016)。赵文智等(2015)针对叠合含油气盆地特点,提出了多勘探“黄金带”观点,认为古老烃源岩“双峰式”生烃、储集层多阶段发育、油气多期成藏和晚期有效,表明了深层油气发现呈多期、多阶段特征。庞雄奇等(2014)将深层油气成藏总结为两种类型,分别为深成油气藏和深埋油气藏,分别是指深部地质条件下形成的油气藏和油气藏在浅部形成后埋藏到深部经改造的油气藏(罗晓容等,2016)。然而,对于这两种深层油气成藏模式,还有多个问题需进一步探讨:①深成型油气藏,油气成藏发生在深部地层,烃源岩热演化程度高,是否还具备有效规模供烃能力?机械压实强烈,储层原生孔隙低,次生孔隙形成机制主要为早期形成-后期保存还是深部形成?能否形成规模有效储集体?深层整体储层致密条件下,油气(尤其是原油)聚集的方式与动力是什么?②深埋型油气藏,油气成藏主要发生在中浅层,烃源岩生烃能力强、储集层尚未达到致密阶段,在成藏要素匹配条件下,油气较容易聚集成藏,但油气藏需经历长期地质演化过程,油气藏封盖条件与后期构造运动是控制油气藏能否保存的关键因素,什么样的保存条件可使油气藏得到保存?

本次研究在调研深层油气成藏要素的基础上,探讨几个与深层油气藏形成及演化相关的科学问

题,提出两种较为可行的深层油气成藏模式,再以塔里木盆地台盆区碳酸盐岩和库车坳陷碎屑岩为例,证实两种深层油气成藏模式的合理性。需要说明的是,塔里木盆地前陆区和台盆区油气成藏模式多样,本次研究重点关注的是现今处于深层—超深层(>4500 m)的油气藏,力图从物质与能量守恒、盆地构造-沉积演化与生排烃史、储层地质流体分析等多个角度,探讨形成现今深层—超深层油气藏的两种较为合理的成藏形式(模式),并非力求涵盖塔里木盆地全部成藏类型。研究成果在深化深层油气成藏理论认识和指导深层油气勘探方面具有一定意义。

1 深层油气藏形成与演化的关键科学问题

关于深层油气成藏环境和成藏要素的分析,前人已开展过详细的论述(庞雄奇等,2014,2020;罗晓容等,2016;刘文汇,2019;操应长等,2022)。普遍认为深部地层具有高温、高压、低孔、低渗地质特征(庞雄奇等,2014)。烃源岩经历了大规模生油期(Tissot and Welte, 1984; Quigley and MacKenzie, 1988),目前主要处于多途径复合生气阶段,生气途径主要包括残余干酪根热裂解、源内残留液态烃热裂解、源外液态烃热裂解和水加氢生气(张水昌等,2021)。深部储层经历了长时间、多地质因素的作用与改造,非均质性强(孙龙德等,2015;罗晓容等,2016),优质储集层发育层位广(古生界至新生界均有发育),与盆地地温梯度有关,“冷盆”内深层油气勘探潜力更大(操应长等,2022)。深层碳酸盐岩储层形成受原始沉积环境和后期成岩改造作用共同控制,台地边缘高能相带礁、滩储集体为后期优质储层形成提供重要基础,断裂活动、白云岩化和热液流体活动对深部优质储层形成尤为关键,斜坡带、不整合面、断裂带附近和流体活动区域成为寻找深部优质碳酸盐岩储层的重要领域(金之钧,2011;何治亮等,2016)。深层碎屑岩优质储层发育是在相对优质的沉积作用基础上,后期受到了次生溶蚀成孔、构造成缝、中—浅层流体超压、早期烃类充注和早期浅埋-晚期快速深埋等一种或多种成岩和地质作用的综合影响下而形成(Bloch et al.,2002;贾承造和庞雄奇,2015;远光辉等,2015;操应长等,2022),具有中—浅层成储、深埋保持和中—浅层成储、深埋调整的两种优质储层发育模式(操应长等,2022)。盖层封盖能力和有效保存条件是深层油气能否成藏的关键(金

之钧,2014;刘文汇,2019)。深层一方面发育多套储盖组合,纵向上相互叠置,表现为多级封盖的特征(何治亮等,2016),另一方面受高温、高压、强应力、多期构造运动影响,盖层力学性质较浅层变化大,碳酸盐岩和泥岩盖层可塑性变差,容易产生裂缝,导致封盖能力变弱,膏盐岩盖层可塑性强,不易破裂,封盖能力强,成为深层最为有效的盖层(金之钧,2010,2011),据统计盖层中仅占8%的膏盐岩盖层,封盖了全球50%以上的油气资源(何治亮等,2016)。

在深层油气藏的形成与演化研究中,还存在一些基本问题需要进一步关注。

1.1 遵循物质守恒和能量守恒基本定律

物质守恒(mass balance)和能量守恒(energy conservation)为自然界普遍存在的两大基本定律,本质是物质和能量都不会凭空产生或创造,也不会凭空消失,它们只能从一种物质转移到另一种物质或从一个物体传递给另一个物体。含油气盆地深层中油气生成、储集空间形成、油气相态和运移等研究均需遵循两大基本定律。

深层存在多种类型烃源,可动态转化接替供烃(赵文智等,2015;刘文汇,2019;张水昌等,2021),干酪根一般在中—浅层大量生成液态烃,此时 R_o 处于0.6%~1.2%的“液态窗”阶段,大量烃源岩热模拟实验表明,I、II型干酪根初次裂解以生油为主,生油量可占总生烃量80%~90%(Espitalié et al., 1988; Burnham and Braun, 1990; Dieckmann et al., 1998),然而初次裂解生成的原油在生烃高峰期的排烃效率普遍小于60%,会随着热解程度升高而逐渐增大(彭平安和贾承造,2021),表明深层烃源岩内可能还存在大量液态烃滞留(Jarvie et al., 2007; 赵文智等,2011),可作为深埋条件下天然气成藏的有效气源灶(赵文智等,2015)。大量热解实验证实,I、II型有机质或干酪根初次裂解生气下限可延至 $R_o=3.5%$ (Mi Jingkui et al., 2018; 彭平安和贾承造,2021),进一步阐明了深层烃源岩残留有机质仍具有一定生气能力,当有机质处于 $2.0% < R_o < 3.5%$ 阶段,生气量占总生气量的20%~30%(张水昌等,2021)。深部外源氢的加入和过渡金属元素催化可在一定程度上提高一过成熟有机质的生气潜力(金之钧等,2002;刘国勇等,2005;张水昌等,2021),但由于深部烃源岩非常致密,深大断裂虽然沟通了深部流体与烃源岩层,但可能只在局部范围内发生外源流体与烃源岩的相互作用,对总的生气量提高多少还有待深入研究。因此,从物质守恒角

度来看,整体上当储层处于中—浅层时,烃源岩供烃能力最强,随着埋深增大、温度升高,生油能力快速降低,形成多途径供气模式,生气规模也会随烃源岩成熟度升高而显著降低。

深层是否依然具备形成规模性储集空间的能力?从物质守恒角度分析,地层自沉积以来至深埋过程中,压实减孔作用一直存在,具有浅层压实减孔作用强烈、中—深层压实作用减弱的趋势(Allen and Allen, 2005),建设性成岩作用普遍发生在中—浅层、破坏性成岩作用多发生在中—深层。受米兰科维奇旋回影响,海平面高频升降(Hinnov, 2013; 吴怀春和房强,2020),沉积期碳酸盐岩频繁地周期性暴露地表,在大气淡水淋滤作用下易于产生溶蚀作用(杨磊磊等,2020)。沉积期—准同生期,在蒸发泵和回流渗透作用下,容易发生规模性白云岩化作用,方解石($CaCO_3$)转换为密度更高的白云石($CaMg(CO_3)_2$),增孔作用发生。表生期,地层长期暴露地表,规模性岩溶作用发生,形成岩溶型孔、洞、缝系统(何治亮等,2016)。中—浅层碎屑岩和碳酸盐岩地层为开放型流体环境的概率更高,长石和方解石等溶蚀产物容易被带走,溶蚀增孔作用显著。此外,中—浅层埋藏阶段若存在早期原油充注,一方面利于深部储层物性的保持(Wilkinson et al., 2006),另一方面可导致储集层矿物润湿性反转,使得部分储集层毛细管力转变为油气运移的“动力”,极大降低深层油气成藏的动力需求,有利于油气运移(Tweheyo et al., 1999)。然而,深层—超深层地层多为封闭型流体环境且地层水矿化度较高,还具备多大溶蚀能力?溶蚀后的产物不易被带出,若发生就近沉淀,一般不会对储集性能产生规模性提升。前人大量研究已证实,深部热液会局部改善碳酸盐岩储集性能(金之钧等,2006;焦存礼等,2011;沈安江等,2015)。热液活动一般发生在深大断裂带,对储集层能否形成规模性改善?在深层—超深层热液溶蚀后的产物去哪了?这些问题还需从物质平衡和能量守恒角度深入探讨。

含油气盆地中油气运移通常被描述为一个或一组未知的、使油气从源到储的过程(<http://wiki.aapg.org/Migration>)。油气二次运移的主要动力是烃-水密度差产生的浮力($\Delta\rho gh$),主要阻力是毛细管力($2\sigma\cos\theta/r$),与孔喉半径、界面张力和岩石润湿性相关(Berg, 1975; Schowalter, 1979)。含油气盆地深层岩石致密,水润湿相储层毛细管阻力大,石油仅在浮力作用下一般难以直接进入致密储层。然

而,在深层高温、高压条件下,天然气中的石油溶解度可大幅提升,经计算地层在埋深 5 km 处,每 1 kg 天然气就可溶解约 1 kg 石油(Batalin and Vafna, 2017)。烃源岩内干酪根或残余液态烃经热裂解形成天然气,压力急剧增大(Guo Xiaowen et al., 2011),超过岩石破裂所需的最小压力,形成微裂隙(Meshcheryakov, 2011),饱和原油的高压气体排出并在浮力作用下二次运移,气体进入深层致密储层中的阻力相对较小,随着气体向上运移或已形成的气藏抬升,油、气相态分离,可形成油气藏。

1.2 注重动态演化过程

含油气盆地深层一般经历了盆地演化的全过程,并最终演化到现今深埋阶段。不同的构造-埋藏演化过程会导致烃源岩生排烃期、储层演化过程和油气成藏期的差异,直接影响深层油气成藏模式的建立,因此需要从动态演化角度关注深层油气成藏。

通过对国内外多个含油气盆地构造-埋藏演化历史的分析与梳理,初步总结了四种盆地构造-埋藏演化模型,分别为:① 长期浅埋、晚期快速深埋模式(图 1a),以塔里木盆地库车坳陷为例(戴金星等, 2012;王招明, 2014);② 长期匀速持续埋藏模式(图 1b),以渤海湾盆地渤中、渤东坳陷为例(侯贵廷和钱祥麟, 1998;漆家福, 2004);③ 早—中期持续深埋、晚期持续抬升模式(图 1c),以鄂尔多斯和四川盆地为例(刘池洋等, 2006;王学军等, 2015);④ 早期快速深埋、后期缓慢埋藏模式(图 1d),以塔里木

盆地顺托果勒低隆起至满加尔凹陷为例(徐国强等, 2005)。在盆地演化时间(250 Ma)、烃源岩厚度(500 m)、有机质丰度(2%)、氢指数(600 mgHC/g TOC)、古今地温梯度(2.5 °C/100 m)和最终埋藏深度(8000 m)设置一致的前提下,开展了四种概念模型的烃源岩生排烃演化过程数值模拟,结果表明(图 2),烃源岩具有快速(持续)埋藏发生越早、主生排烃时期越早的趋势;早期快速埋藏,烃源岩成熟快,生排烃时间早,后期烃类生成潜力急剧降低;长期浅埋,烃源岩成熟度低,生烃转换率低,晚期快速深埋,烃源岩快速生排烃;Ⅲ型有机质的主生排烃时期滞后于Ⅰ、Ⅱ型有机质。从动态演化角度分析,中—浅层储集层物性较好,当与烃源岩主生排烃时期匹配时,油气最可能聚集成藏,早期快速(持续)深埋演化模式最有可能对应于油气早期成藏、后期深埋保存的成藏模式,保存条件是关键,而长期浅埋—晚期快速深埋演化模式最可能形成深层—超深层(凝析)油气聚集成藏模式。

1.3 探索深层液态烃稳定性与保存下限

近年来,随着油气勘探的不断深入,国内外深层—超深层领域发现了一系列高温油藏,如北海盆地 Elgin, Franklin 油田在 185~203 °C 的三叠系和侏罗系砂岩储层中发现凝析油藏(Pepper and Dodd, 1995; Waples, 2000),冀中坳陷在 5642~6027 m 埋深、井底温度为 201 °C 的中元古界雾迷山组发现凝析油藏(赵贤正等, 2011),塔里木盆地轮

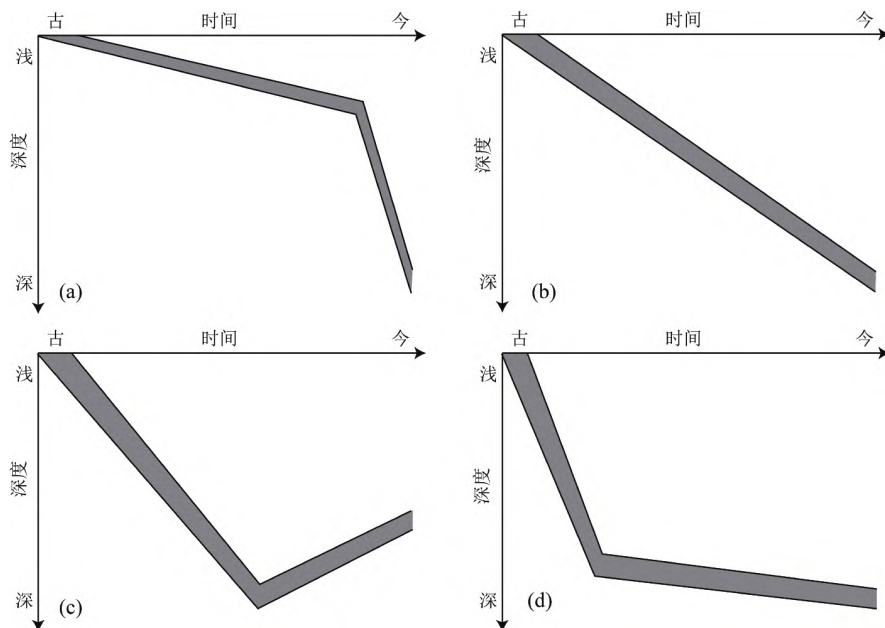


图 1 含油气盆地不同类型构造-埋藏演化模式

Fig. 1 Various tectonic evolution and burial history models of petroliferous basins

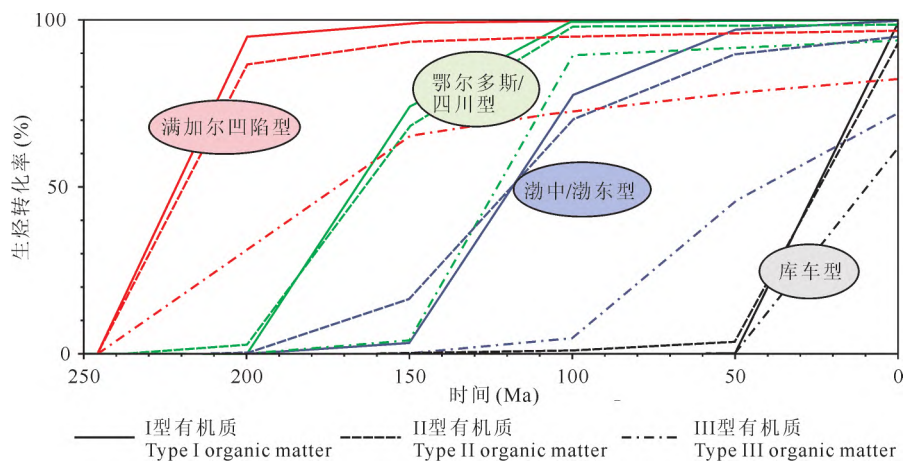


图2 不同构造-埋藏演化条件下的烃源岩生烃转换率随时间演化曲线

Fig. 2 Transformation ratios of hydrocarbon generations versus evolution timing under different tectonic and burial conditions

探1并在8203~8260 m埋深、井底温度为161℃的寒武系储层中获得轻质原油(杨海军等, 2020),塔深1并在8404~8406 m的寒武系白云岩储层钻遇褐黄色液态原油(翟晓先等, 2007)。这些勘探成果突破了传统的“液态窗”和勘探“黄金带”理论,表明原油的稳定性比早期推测的结果要高很多,在超过160℃的高温环境中原油仍然可以稳定存在。

前人基于原油热裂解生烃模拟实验及化学动力学软件开展了大量原油稳定性及保存温度上限的研究(Waples, 2000; 田辉等, 2006; Zhu Guangyou et al., 2012)。事实上,实际的地质条件往往极其复杂,除了温度这一重要因素外,仍有多因素影响原油裂解成气的重要因素未被认知。一些学者认为高压能促使原油发生裂解,并加速油气的生成与演化过程(Braun and Burnham, 1990),而以Dominé and Enguehard (1992)、Behar and Vandenbroucke (1996)和Hill et al. (1996)为代表的学者研究认为,超压环境会抑制原油裂解,同时,裂解产物中轻烃组分含量降低、重烃组分含量增加。对于水在原油裂解过程中的影响,部分学者认为水的存在能提高热解体系中原油的稳定性,导致原油裂解气的产量大幅降低(Behar et al., 2003)。另有学者认为水可以作为溶剂与油中组分形成含水有机相,并随着温度升高可以为原油裂解提供额外的氢源(Lewan, 1992; Seewald, 1994; Seewald et al., 1998)。此外,帅燕华等(2012)基于黄金管热模拟实验提出地层水中 Mg^{2+} 对原油+水的反应有一定促进作用,使烷烃气总量、 H_2 与 CO_2 产量均有所增加。不同的介质条件对原油裂解能够产生一定影响(赵文智等, 2007);碳酸盐岩可以显著降低甲烷生成的活化

能,引起原油裂解温度的降低,其次是泥岩,而砂岩影响最小。矿物的催化对原油稳定性及原油裂解气的生成也起着非常关键的作用。一般认为,蒙脱石可以大大促进原油催化裂解成气(Tannenbaum and Kaplan, 1985),而伊利石则被认为有利于原油转化为凝析油或者天然气(Espitalie et al., 1980)。TSR作用能够促使原油在较低的温度下快速裂解成气,同时产生大量 H_2S (Seewald, 2003)。因此,关于影响深层液态烃稳定性的因素,目前还存在较大争议,而且常用的封闭体系黄金管热模拟实验方法,无法实现对裂解原油相态与压力的原位、可视化定量分析,需要探索新的实验方法(如人工合成流体包裹体技术),加大对深层液态烃稳定性与保存深度下限的研究,以期在丰富原油裂解成气理论和指导深层勘探决策(找油或找气)方面发挥重要作用。

1.4 融合多学科与多技术

深层油气资源总体上具有地层年代老、构造改造强、温度-压力高、储层类型和流体相态多样的特征,经典流体力学、表面化学、化学动力学、分子动力学和物理与数值模拟等多学科交叉是解决深层油气藏形成与演化问题的必然趋势,其中分子动力学模拟是连接宏观与微观、实验与理论的桥梁。

深层埋藏条件复杂,现有技术无法测试地质条件下高温高压及各向应力异性,无法反映原位应力下的微纳米裂缝,导致深层碎屑岩储层评价困难,亟需发展真实地质条件下的储层物性测试装置及数值模拟技术。

目前的有限(<3)变量的物理和数值模拟都无法完全呈现深层油气所经历的复杂地质过程,因此亟需开展真实地质条件下油气演化全过程研究。盆

地和含油气系统模拟技术为仿真研究深层油气演化提供了一个定量和半定量的研究方法,但目前还存在一些技术瓶颈(贾承造等,2021),例如,如何解决热-流-固-化学(THMC)多场耦合在盆地尺度中的应用,在分子动力学和生烃动力学研究方面,如何链接对微观行为认知和油气在盆地尺度的宏观响应,如何考虑真实地质条件下全过程生、排、运、聚和改造。这需要在盆地中开展全油气系统的研究(贾承造,2019)。

2 塔里木盆地深层油气成藏模式

综合上述分析认为,深层油气藏的形成可能主要为以下两种类型成藏模式:① 中—浅层成藏、深埋保持型;② 长期浅埋、晚期快速深埋成藏型。下面以塔里木盆地台盆区顺北地区和库车坳陷博孜—大北地区为例开展论述。

2.1 区域地质概况

塔里木盆地位于新疆维吾尔自治区境内,是中国最大的含油气盆地,总面积约为 $56 \times 10^4 \text{ km}^2$ (图 3; Jia Chengzao and Wei Guoqi, 2002)。该盆地属于典型的叠合盆地,由古生界海相克拉通盆地与中、新生代陆相前陆盆地构成,在经历了多旋回构造演

化阶段后,形成了现今的构造格局(Jia Chengzao and Wei Guoqi, 2002)。塔里木盆地的地层沉积充填序列可以简要概括为震旦系—泥盆系海相沉积体系、石炭系—二叠系海陆过渡相沉积体系以及中生界—新生界陆相沉积体系。盆地内油气资源量巨大,常规石油资源量为 $75.06 \times 10^8 \text{ t}$,天然气资源量为 $12.94 \times 10^{12} \text{ m}^3$,油气当量超过 $178 \times 10^8 \text{ t}$,且主要分布在深层和超深层领域(杨学文等,2021)。目前塔里木盆地深层—超深层油气勘探主要集中在台盆区碳酸盐岩和库车坳陷碎屑岩领域,其中台盆区目的层系为奥陶系、寒武系和震旦系,库车坳陷深层目的层主要是白垩系和侏罗系(图 4)。

2.2 台盆区顺北地区深层油气成藏模式

前人对塔里木盆地台盆区顺北—塔北地区奥陶系油气成藏开展了大量研究,主要形成以下四种不同观点:① 海西晚期(约 250 Ma)一期油气成藏模式(Zhu Guangyou et al., 2012, 2019; Ge Xiang et al., 2020; Li Jingfei et al., 2020);② 加里东晚期(约 436~405 Ma)和喜马拉雅晚期(约 20~2 Ma)两期油气成藏模式(Gong et al., 2007; Wang Tieguan et al., 2008; Fang Ronghui et al., 2017; Li Meijun et al., 2018);③ 海西中—晚期(约 327~250 Ma)和喜马拉雅晚期(约 6~2 Ma)两期油气成

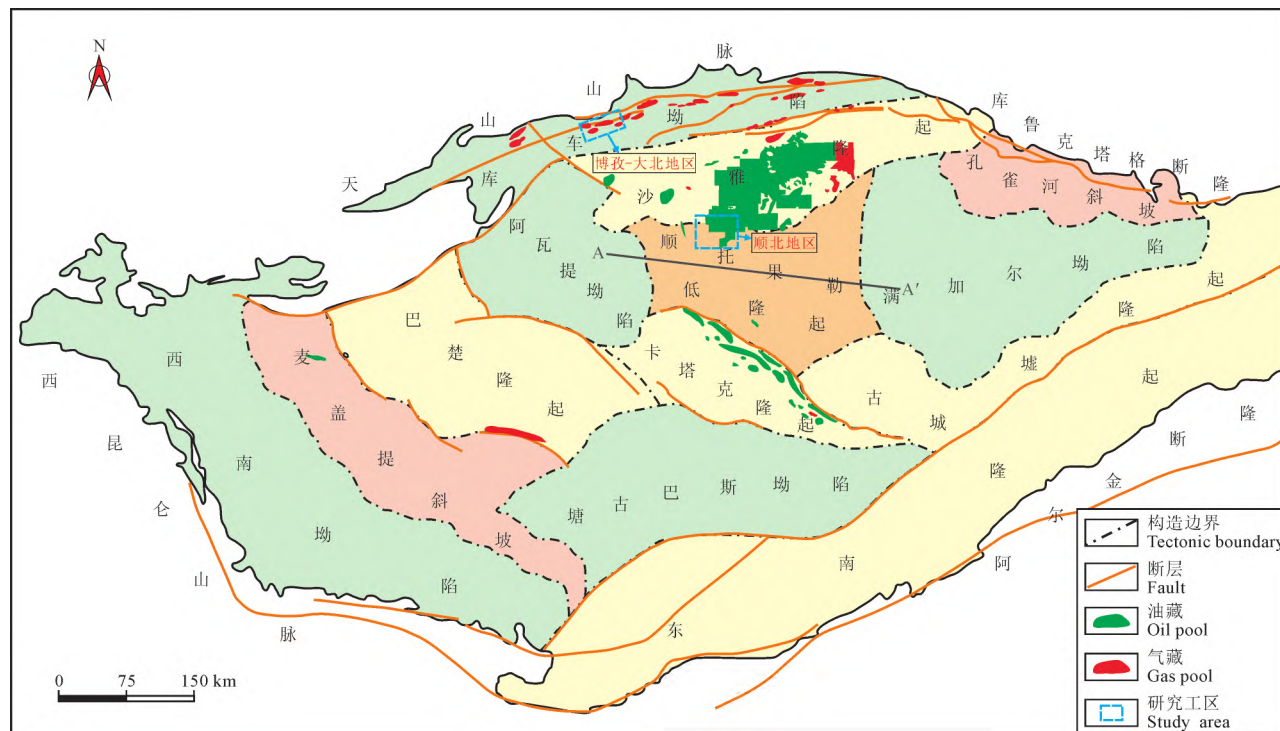


图 3 塔里木盆地油气藏分布图及研究工区位置(据李阳等,2020 修改)

Fig. 3 Distribution of oil and gas reservoirs in the Tarim basin and location of the study areas shown in blue dash boxes (modified from Li Yang et al., 2020)

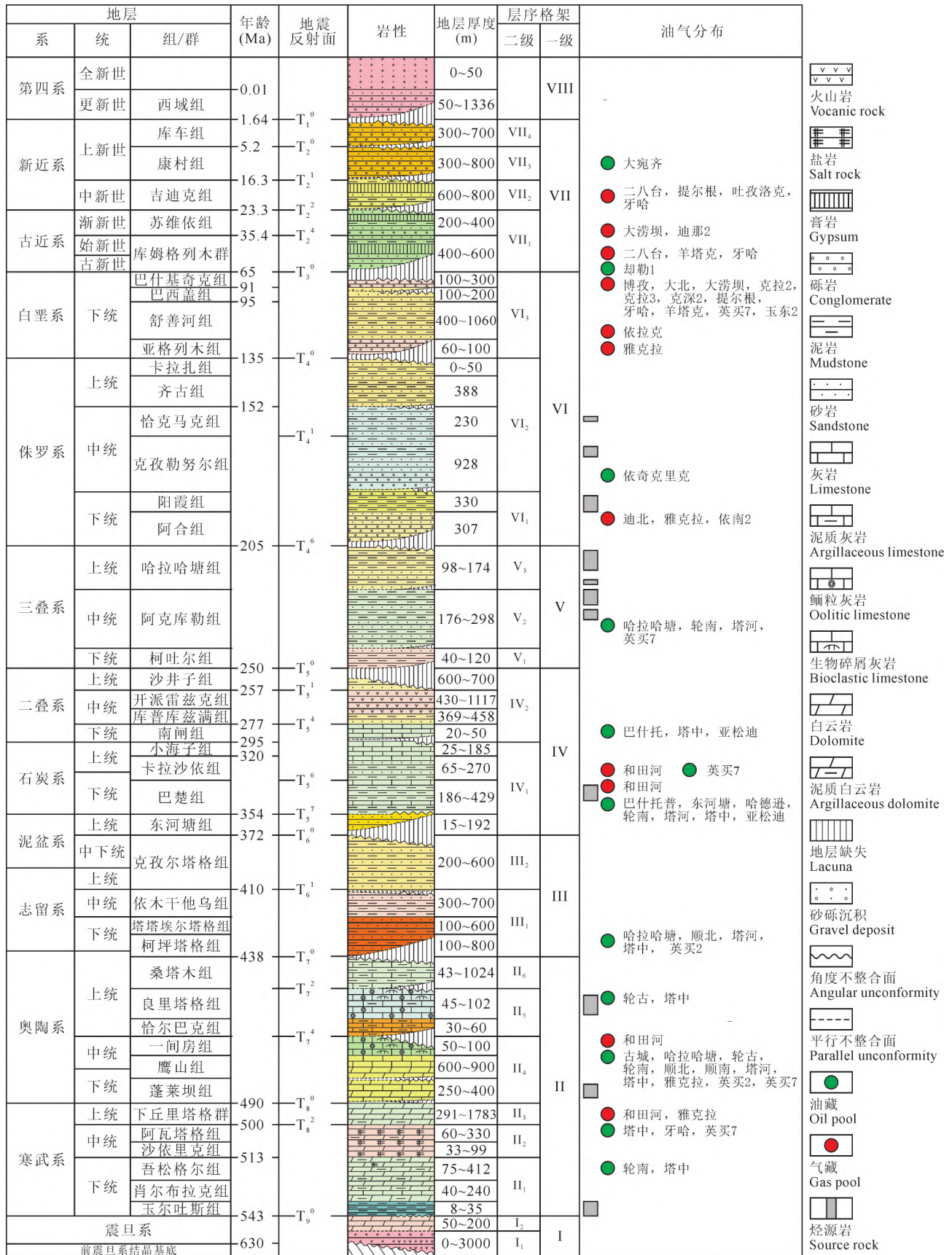


图4 塔里木盆地地层综合柱状图 (据 Lin Changsong et al., 2012 修改)

Fig 4 Generalized stratigraphic column of the Tarim basin (modified from Lin Changsong et al., 2012)

藏模式(He Dengfa et al., 2002; 饶丹等, 2014); ④ 加里东中—晚期至海西早期(约 463. 2~376 Ma)、海西晚期(约 312. 9~255. 0 Ma)、燕山期(约 150. 2~100. 6 Ma)和/或喜马拉雅期(约 22. 0~2. 0 Ma)三~四期油气成藏模式(陈红汉等, 2014; 王玉伟等, 2019; 顾忆等, 2020)。这些研究普遍是基于流体包裹体系列分析与盆地模拟相结合得到, 但研究过程中一些关键参数(如, 流体包裹体测温值、古热流值和地层剥蚀量等)获取的准确与否, 直接影响着油气成藏年代厘定的可靠性, 尤其类似塔里木盆地这样构造演化复杂、地层古老、埋藏深的盆地, 准确获取这些关键参数难度极大, 这可能是导致该地区深层油气成藏模式多样的重要因素。

近年来, 方解石激光原位 U-Pb 测年技术发展快速(Coogan et al., 2016; Roberts et al., 2016, 2020)。与常规的同位素稀释法相比, 该技术具有样品制备简单、空间分辨率高和测量效率高等优点, 能够实现对多世代碳酸盐胶结物的原位精确测年(Godeau et al., 2018), 并已成功应用于厘定洋壳中碳酸盐矿物的形成年龄(Coogan et al., 2016)、刻画断层活动事件(Roberts et al., 2016; Nuriel et al., 2017)以及揭示碳酸盐岩储层的流体演化历史等研究(Godeau et al., 2018; Yang Peng et al., 2022b)。特别指出的是, 当方解石的胶结作用与石油充注事件同时发生时, 储层中的原油可以在方解石的晶格缺陷中被捕获, 形成原生流体(油)包裹体(Goldstein et al., 1994)或者经过后期蚀变作用形成与方解石胶结物共生的固体沥青。这种情况下, 方解石胶结物的形成时间便可近似代表石油充注发

生的时间(Rochelle-Bates et al., 2021; Yang Peng et al., 2022b), 为厘定具有复杂构造演化历史的古老深埋沉积盆地中的石油充注历史提供了新的思路。

本次研究综合方解石激光原位 U-Pb 测年技术、石油和碳酸盐岩储层地球化学分析、流体包裹体分析以及盆地模拟等方法, 重建顺北地区奥陶系深层海相碳酸盐岩油气充注历史, 建立台盆区深层油气成藏模式。

2.2.1 原油地球化学特征及类型

根据原油获取方式的不同, 首先将原油分为储层产出原油、岩样萃取原油和包裹体分离原油, 分别是指奥陶系现今产层中的油、碳酸盐岩岩石粗碎后萃取得到的油和碳酸盐岩包裹体破碎后抽提得到的油。分别对三种原油开展油气地球化学分析, 结果表明, 姥鲛烷/ nC_{17} (Pr/nC_{17}) 与植烷/ nC_{18} (Ph/nC_{18}) 比值分别分布在 0.15~0.79 和 0.30~0.84, 根据 Connan and Cassou (1980) 评价标准, 认为三种原油均来源于海相、II 型干酪根的烃源岩(图 5a)。此外, 顺北地区奥陶系储层原油、游离油和包裹体油的甾烷、萜烷及其他地化特征也较为接近(Yang Peng et al., 2021), 表明三种原油可划分为同一油族, 推测对应的烃源岩形成于高度还原的海相环境, 有机物质主要来自于藻类(特别是绿藻)和细菌。原油中甲基化芳烃的异构化比值常被用作指示原油热成熟度(Alexander et al., 1985; Radke et al., 1994), 根据该指标, 可将三种原油进一步划分成两种类型: A 类原油的甲基菲指数 MPI-1 主要分布在 0.71~0.88 之间, 根据换算关系(Radke,

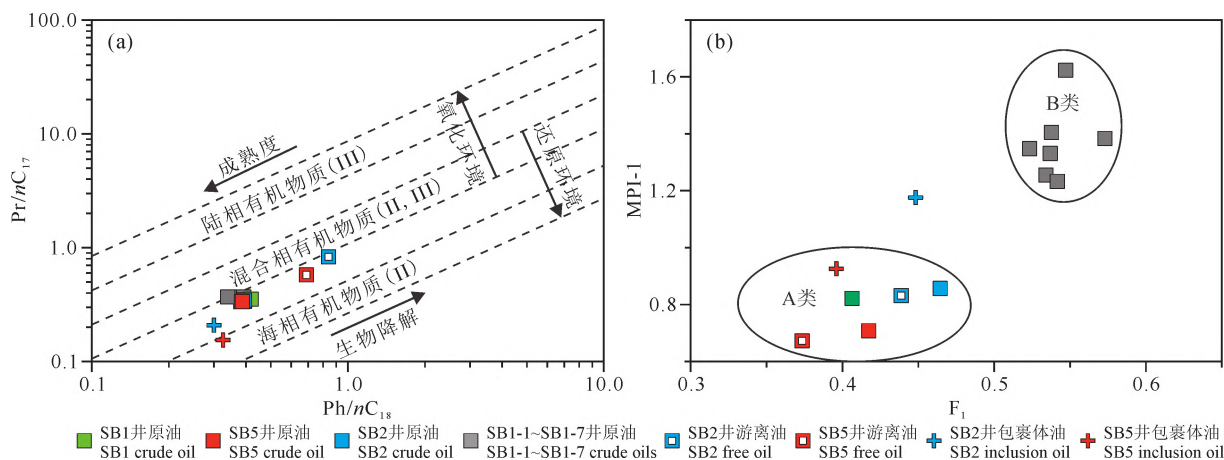


图 5 顺北地区储层原油、游离油以及包裹体油 Ph/nC_{18} 与 Pr/nC_{17} (a) 以及 F_1 与 MPI-1 (b) 比值交会图

Fig 5 Cross plots of Ph/nC_{18} versus Pr/nC_{17} (a) and methylphenanthrene index MPI-1 versus methylphenanthrene ratio F_1 (b) for reservoir oils, free oils and inclusion oils in the Shunbei area

1988; Boreham et al., 1988), 计算得到的烃源岩等效镜质组反射率值 R_c 在 0.82%~0.96% 范围内, 处于生油早期阶段; B 类原油的 MPI-1 值为 1.23~1.40, 计算得到的烃源岩等效镜质组反射率值 R_c 在 1.14%~1.24% 之间, 处于生油晚期阶段(图 5b)。

2.2.2 流体包裹体类型与显微测温

根据油包裹体的气-液比、荧光颜色及沥青含量等方面差异, 可将顺北地区奥陶系油包裹体划分为

两种类型。I 类油包裹体组合具有荧光下为乳白色、气体比约为 5%、常温下呈气-液两相(L_{oil} -V)或者气-液-固三相(S_{bit} - L_{oil} -V)的特征(图 6a~d)。大部分 I 类油包裹体沿方解石胶结物中的愈合裂隙分布, 部分沿方解石胶结物的生长带集中分布(图 6e、f)或者在方解石晶体内部零散状分布, 表现出原生流体包裹体的典型特征。II 类油包裹体组合具有发亮蓝色荧光、气体充填度一般小于 5%、常温下呈

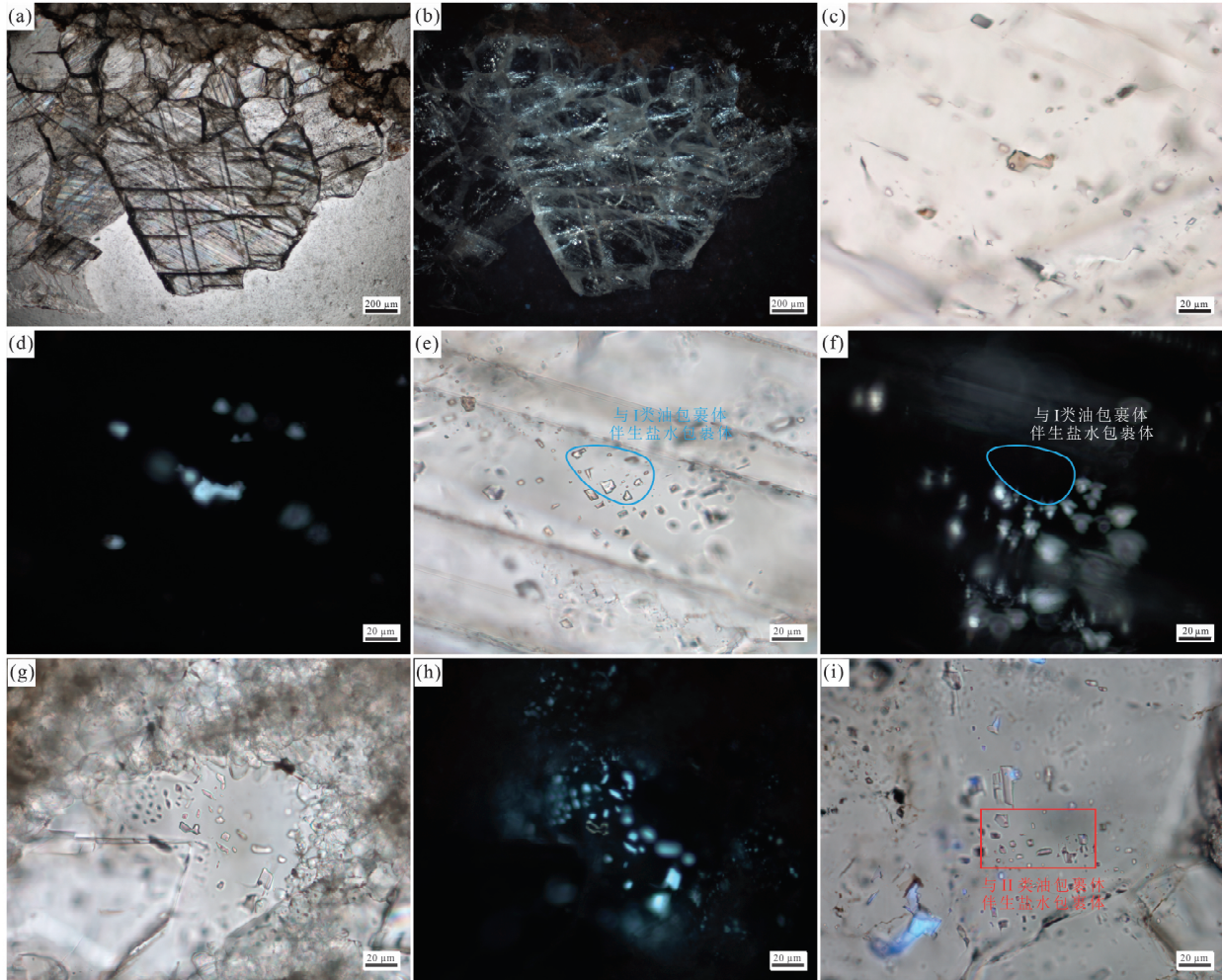


图 6 顺北地区奥陶系储层流体包裹体单偏光与荧光显微照片

Fig. 6 Paired photomicrographs of fluid inclusions under plain transmitted light and UV illumination in the Ordovician reservoirs in the Shunbei area

(a, b)—I 类油包裹体组合主要出现在方解石胶结物的愈合微裂隙中, SB5 井, 7427.30 m; (c, d)—发近白色荧光的、气-液-固三相(S_{bit} - L_{oil} -V)的 I 类油包裹体, 固体沥青沿包裹体壁发育, SB5 井, 7427.30 m; (e, f)—I 类油包裹体与伴生的盐水包裹体沿着 C2 方解石胶结物的生长带分布, SB5 井, 7427.30 m; (g, h)—发近蓝色荧光的两相(L_{oil} -V)油包裹体(II 类油包裹体组合), SB2 井, 7361.50 m; (i)—II 类油包裹体与伴生的盐水包裹体, SB2 井, 7361.50 m

(a, b)—group I oil inclusions assemblages occurring mainly along annealed microfractures within the calcite cements, well SB5, 7427.30 m; (c, d)—near white fluorescing, triphasic (S_{bit} - L_{oil} -V) group I oil inclusions at room temperature, with solid bitumen occurring along inclusion walls, well SB5, 7427.30 m; (e, f)—group I oil inclusions and coeval aqueous inclusions occurring along the growth zone of the C2 calcite cement, well SB5, 7427.30 m; (g, h)—near blue fluorescing, diphasic (L_{oil} -V) group II oil inclusions at room temperature, well SB2, 7361.50 m; (i)—group II oil and coeval brine inclusions occurring along annealed microfractures, well SB2, 7361.50 m

液-固两相 ($S_{bit}-L_{oil}$) 或者气-液-固三相 ($S_{bit}-L_{oil}-V$) 的特征,一般沿切割方解石晶体的愈合裂隙发育(图 6g~i)。

两种类型油包裹体在显微荧光光谱和红绿焰 ($Q_{650/500}$) 方面差异明显(图 7a): I 类油包裹体的最大荧光强度对应波长 (λ_{max}) 和 $Q_{650/500}$ 分别在 533~541 nm 和 0.69~1.03(平均值:0.85); II 类油包裹体的 λ_{max} 和 $Q_{650/500}$ 分别为 522~530 nm 和 0.33~0.54(平均值:0.40)。与此同时,选取能代表两种类型原油的储层原油进行荧光光谱分析,结果表明,A 类原油(SB 1 井和 SB 5 井原油)的显微荧光光谱 λ_{max} 和 $Q_{650/500}$ 值均较大,整体与 I 类油包裹体的荧光光谱特征较为相似;B 类原油(SB 1-1 井和 SB 1-3 井原油)的显微荧光光谱 λ_{max} 和 $Q_{650/500}$ 值相对较低,与 II 类油包裹体的荧光光谱特征吻合度较高(图 7a)。综合油包裹体荧光光谱特征与原油地化特征认为,A 类原油与 I 类油包裹体相似,成熟度相对较低;B 类原油与 II 类油包裹体相似,成熟度相对较高。

顺北 5 井奥陶系碳酸盐岩储层中均可观测到两种类型油包裹体,对油包裹体及其伴生的盐水包裹体均一温度 (T_h) 和盐度测试,结果表明(图 7b): I 类油包裹体 T_h 值分布在 44.5~74.2°C 之间,与其伴生的盐水包裹体 T_h 值在 77.5~95.3°C 之间,冰点温度 (T_m) 值分布在 -5.2~-0.6°C 之间,计算得到盐度值为 1.05%~8.14%NaCleq(平均值:4.18%NaCleq); II 类油包裹体 T_h 值在 39.9~59.1°C 之间,伴生的盐水包裹体 T_h 值为 93.5~126.2°C, T_m 值分布在 -15.2~-8.2°C 之间,计算得到盐度值为 11.93%~18.80%NaCleq(平均值:

15.59%NaCleq),也表现出两种类型油包裹体的明显差异。

2.2.3 原油充注时期厘定

顺北地区奥陶系储层可识别出两期方解石胶结物,分别定为 C1 和 C2,其中 C1 胶结物作为裂缝充填物产出,多为中-粗晶方解石,呈块状或者马赛克状集合体,发暗红色阴极光或者不发光;C2 方解石胶结物同样以裂缝充填物的形式出现并切割 C1 方解石胶结物,主要由发橘黄色阴极光的块状或者马赛克状方解石集合体构成(图 8a、b)。针对两期方解石胶结物,在澳大利亚科廷大学开展原位方解石 U-Pb 定年分析,结果表明:C1 方解石胶结物的 U-Pb 同位素年龄为 446.1±4.4 Ma(1 σ)(图 8c),C2 方解石胶结物的 U-Pb 同位素年龄为 425.7±14.0 Ma(1 σ)(图 8d)。在原油包裹体岩相学观察过程中发现,部分 I 类油包裹体沿着 C2 方解石胶结物生长带集中发育(图 6e、f),表明 I 类油包裹体与 C2 方解石同期形成,可用 C2 方解石的 U-Pb 同位素年龄代表 I 类原油捕获的绝对年龄。

与此同时,利用包裹体均一温度与单井埋藏史-热史匹配,对两类油包裹体的充注时期进行厘定。考虑到方解石中流体包裹体发生再平衡作用的潜在效应(Bourdet et al., 2008),本次研究选择与油包裹体伴生盐水包裹体连续分布均一温度的最小值作为该类油包裹体的捕获温度。结果表明:顺北地区奥陶系碳酸盐岩储层中 I 类油包裹体在约 426 Ma 的早泥盆世(加里东晚期)被捕获,捕获时储层深度在 2000 m 左右,捕获时间与 C2 方解石 U-Pb 定年得到的 I 类原油充注年龄相互验证;II 类油包裹体的捕获时期约为 330 Ma(海西中期),此时储层位于约

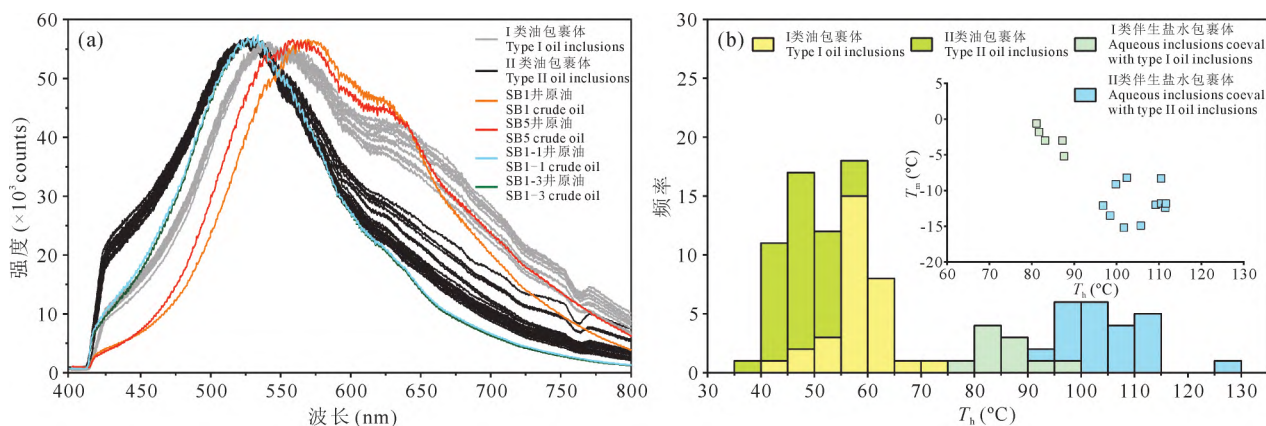


图 7 顺北地区流体包裹体显微荧光光谱(a)与显微测温(b)特征

Fig. 7 Hydrocarbon fluid inclusion fluorescence spectral signatures (a), and microthermometric data (T_h and T_m) (b) of oil and coeval brine inclusions in the Shunbei area

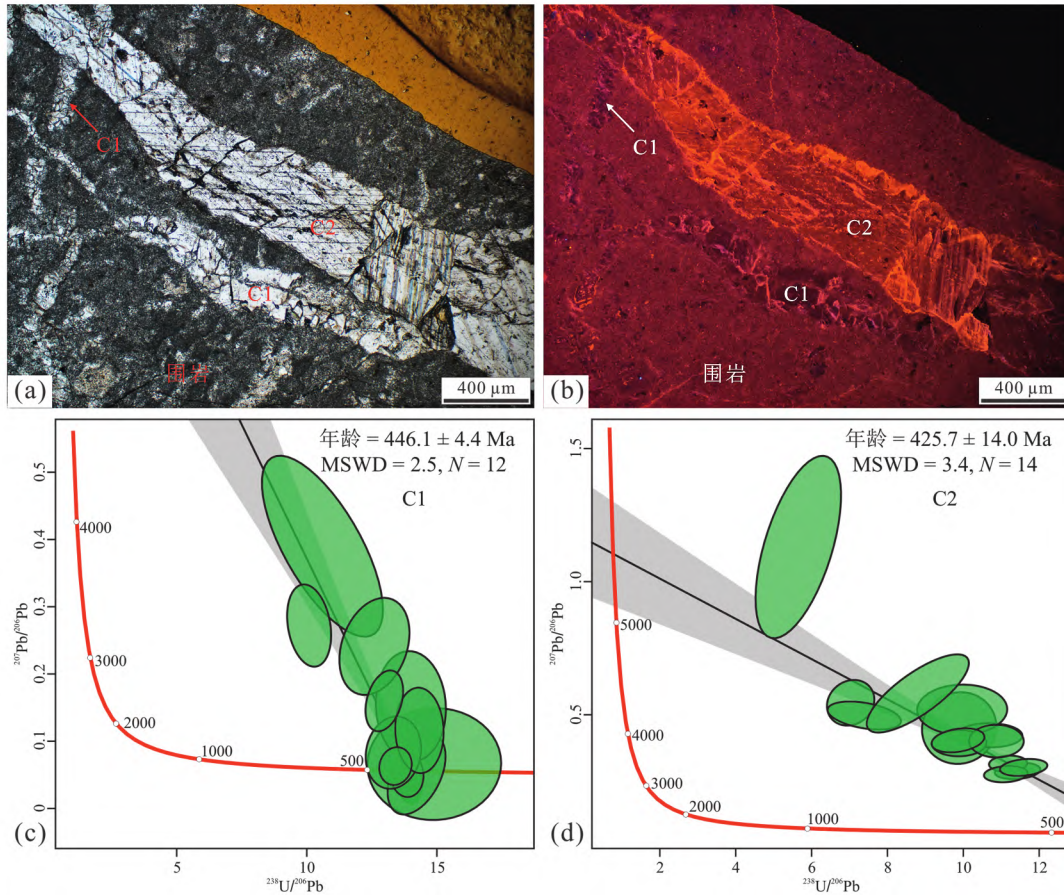


图 8 顺北地区奥陶系碳酸盐岩储层方解石胶结物岩相学特征(a,b)以及 U-Pb Tera-Wasserburg 谐和曲线(c,d)(据杨鹏等,2022a 修改)

Fig 8 Petrographic characteristics (a, b) and U-Pb Tera-Wasserburg concordia plots (c, d) for the calcite cements in the Ordovician reservoir in the Shunbei area (modified from Yang Peng et al. ,2022a)

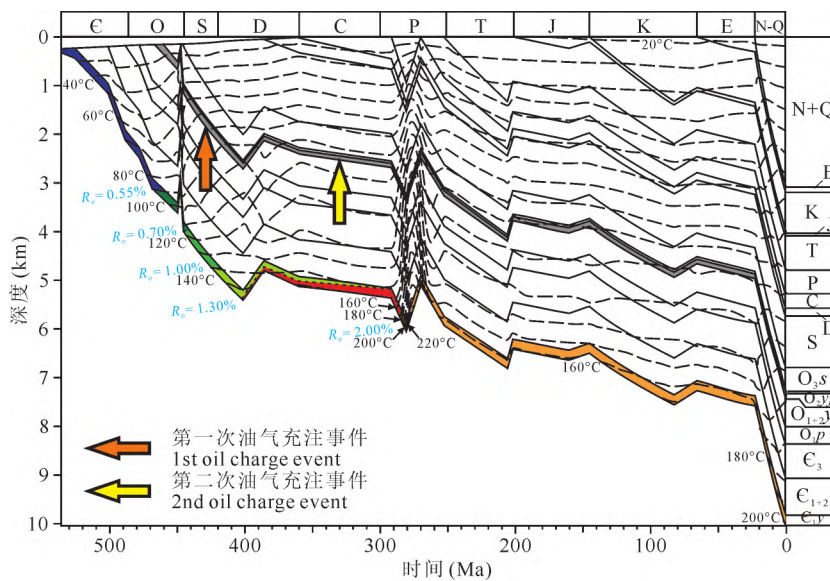


图 9 顺北地区顺北 5 井埋藏史-热演化史及油气充注时期

Fig 9 Burial-thermal history and hydrocarbon charge timing of well SB5 in the Shunbei area

2600 m 埋深(图 9)。由此可见,顺北地区奥陶系现今的深层油气藏均为相对早期的中—浅层成藏。

2.2.4 油气成藏过程与模式

(1)第一期低成熟度原油充注(加里东晚期):加里东早期(寒武纪—早奥陶世),顺托果勒地区广泛发育稳定的台地相碳酸盐岩,在与沉积间断相关的不整合和大气淡水岩溶作用影响下,形成广泛的岩溶型储层。加里东中期(奥陶纪),受到卡塔克隆起与沙雅隆起整体隆升的影响,顺托果勒地区形成近

南北向的构造低隆带,并广泛发育断穿寒武系的北东向和北西向走滑断层。晚奥陶世,伴随桑塔木组巨厚泥岩的沉积,顺托果勒地区奥陶系形成完整的生-储-盖组合(图 10a)。加里东晚期(志留纪),顺托果勒地区下寒武统玉尔吐斯组烃源岩成熟生油,石油沿近于直立的走滑断层向上运移,在早泥盆世(约 426 Ma)充注到奥陶系和志留系圈闭,形成大规模油藏(图 10b)。加里东晚期—海西早期(志留纪—泥盆纪),塔中地区冲断与走滑构造变形作用进一步

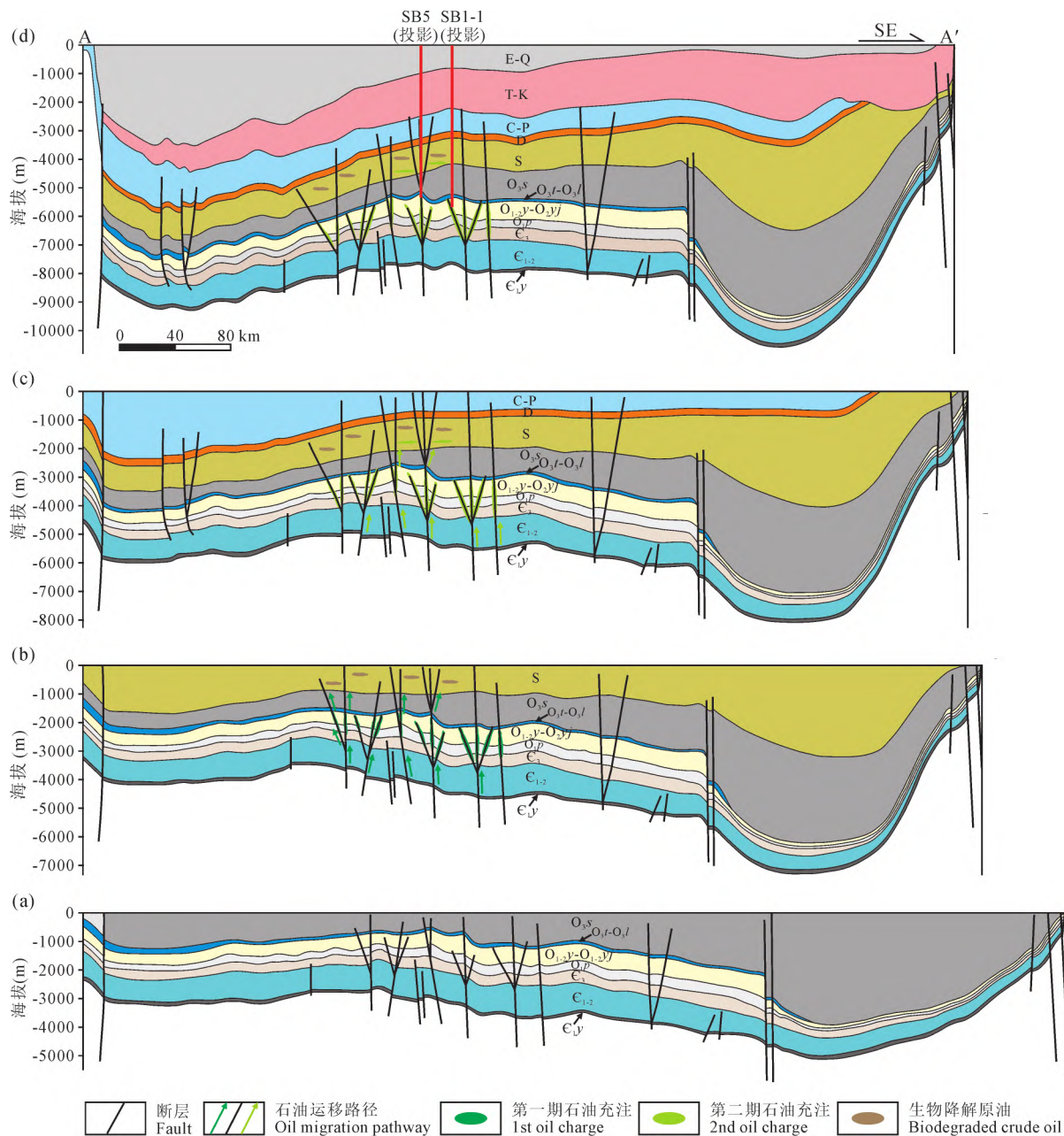


图 10 顺北地区奥陶系深层油气藏成藏演化模式(剖面位置见图 3;构造演化模式据漆立新,2016 修改)

Fig 10 Oil accumulation and charge evolution model for the Shunbei deep oil reservoirs (see Fig 3 for cross section location; the structural evolution model was modified from Qi Lixin, 2016)

强化,顺托果勒地区整体抬升,并继承发育一系列大型走滑断层带,向上扩展至中—下泥盆统。强烈的构造隆升导致顺托果勒地区志留纪地层大规模剥蚀。志留系油藏盖层被破坏,发生严重的生物降解作用,形成广发发育的志留系沥青砂岩。相比之下,下覆的奥陶系油藏由于埋深较大且保存条件较好,并未受到生物降解作用的影响。

(2)第二期高成熟度原油充注(海西中期):海西中—晚期(石炭纪—二叠纪),顺托果勒地区由拉张型应力背景转为挤压型应力背景,发育挤压性或者压扭性断裂。受持续埋藏作用影响,下寒武统玉尔吐斯组烃源岩生成大量高成熟油气,沿着通源走滑断裂向上运移,在中石炭世(约330 Ma)再次充注到顺北地区奥陶系储层(图10c),该时期原油充注时储层中存在超压。由于走滑断裂带的活动时间、强度以及断裂带内部连通性的差异,不同原油充注事件对不同断裂带甚至同一断裂带内不同部位的贡献程度不均,导致顺北地区现今不同断裂带中原油性质的差异。

(3)早期中—浅层形成的油藏持续保存(印支期至喜马拉雅期):印支期—燕山期(三叠纪—白垩纪),顺托果勒地区整体持续沉降,后期多幕构造运动并未对顺北地区产生较大的地质影响,直至晚喜马拉雅期(新近纪—第四纪),才形成现今的构造格局(图10d)。顺北地区奥陶系现今油藏产层中部的地层温度在148~167℃附近,为该油层所经历的最高地层温度,表明原油未发生显著的热裂解作用。储层的地层水为CaCl₂型,表示该地区油藏现今的封闭性良好,有利于油藏保存。

综上所述认为,顺北地区奥陶系深层油气藏为“早期中—浅层成藏、后期深埋持续保存”的成藏模式,油气成藏后相对稳定的构造背景是油气藏能保持至今的关键因素。

2.3 库车坳陷博孜一大北地区深层油气成藏模式

库车坳陷博孜一大北地区深层—超深层碎屑岩天然气资源丰富,目前已在白垩系储集层中勘探发现多个天然气藏(如,博孜9、博孜12、大北9、大北12等气藏)(杨海军等,2019;田军等,2020;王珂等,2022)。前人对大北地区油气成藏开展了一些研究,形成两种类型成藏模式:①库车期以来的深层—超深层两期油气成藏模式(王招明等,2014;Zhao Shuangfeng et al.,2019);②两期原油、一期天然气充注成藏模式,其中第二期原油和晚期天然气充注发生在库车期,为深层—超深层成藏(Guo Xiaowen

et al.,2016;鲁雪松等,2016)。虽然前人针对博孜地区的沉积—储层特征开展了相关研究(曾庆鲁等,2020;王珂等,2022),但对该地区的油气成藏时期、过程与模式研究相对薄弱,加强博孜地区油气成藏研究对深化克拉苏构造带西部白垩系深层油气成藏模式和下一步勘探选区具有一定理论和实际意义。

在对博孜地区白垩系多个天然气藏解剖的基础上,本次研究重点选取博孜102井开展分析,首先对下白垩统巴什基奇克组超深层碎屑岩储层进行取样(6758.5~6866.7 m),然后开展流体包裹体岩相学、光谱学、显微测温以及捕获时古温度、古压力模拟(PVT模拟),得到更符合实际地质的成藏时期古温度、压力数据,在此基础上,结合单井埋藏史—热史开展油气成藏期次与时期分析,明确博孜地区深层油气成藏模式。

2.3.1 流体包裹体发育特征与显微测温

博孜地区下白垩统烃类包裹体主要发育在碎屑颗粒边缘、石英愈合裂隙和胶结物中,部分可见捕获于石英次生加大边中,常以带状、群状或孤立状分布于宿主矿物中,大小分布在3~20 μm之间,大部分集中在5~8 μm之间,包裹体气液比集中在2%~20%之间。根据油气包裹体的产出位置、荧光颜色等特征,可将博孜102井烃类包裹体划分成两类:第Ⅰ类烃包裹体组合主要由气—液两相(L_{oil}-V)油包裹体组成,呈亮白色、蓝绿色荧光,发育在石英次生加大边与碎屑颗粒之间的“尘线”处以及石英颗粒的愈合裂隙中(图11a~d);第Ⅱ类烃包裹体组合主要由气—液—固三相(S_{bit}-L_{oil}-V)油包裹体和沥青(S_{bit})包裹体组成,发育在石英颗粒愈合裂隙内、或沿切穿石英次生加大边的愈合裂隙分布(图11e~h)。

包裹体显微荧光光谱实验结果表明,两种类型油包裹体的荧光光谱特征差异明显,其中第Ⅰ类油包裹体的荧光光谱主峰波长λ_{max}值分布较为集中,在537~540 nm之间,第Ⅱ类油包裹体的λ_{max}值分布跨度较大,在521~575 nm之间,且部分包裹体光谱主峰出现了明显的“蓝移”或“红移”次峰(图11i、j),反映了两种类型油包裹体具有不同的成熟度。第Ⅱ类油包裹体荧光光谱的红绿商Q_{650/500}值跨度较宽、光谱出现明显双峰(535~545 nm、560~567 nm)、气液比变化幅度较大(1.87%~16.82%)以及观察到许多气—液—固三相或沥青包裹体等特征,均表明第Ⅱ类油包裹体是早期充注原油受到了后期天然气气洗改造而捕获形成,代表了天然气充注事件。

流体包裹体显微测温结果表明,第Ⅰ类油包

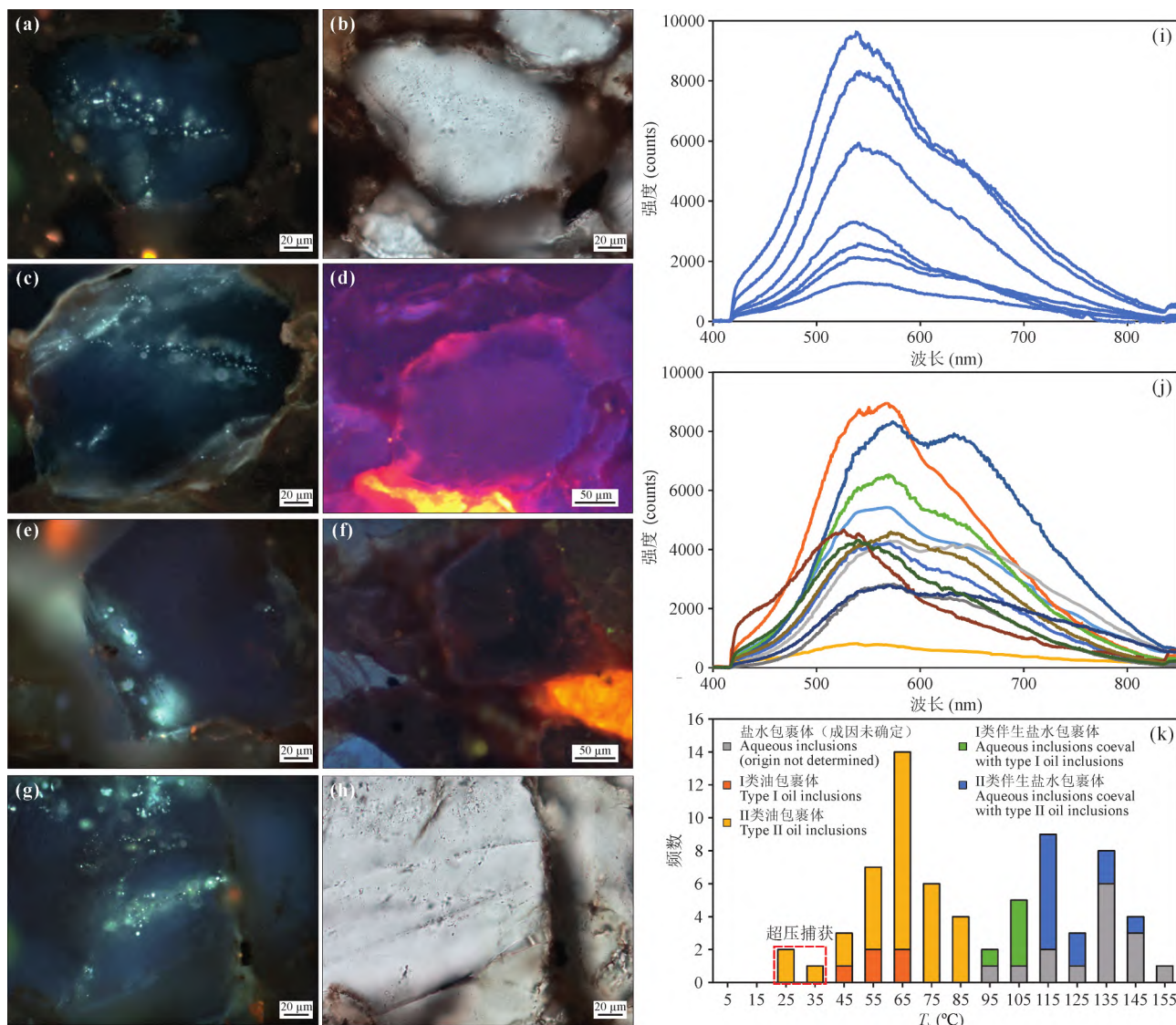


图 11 博孜 102 井白垩系烃类包裹体岩相学、显微荧光特征和流体包裹体均一温度

Fig. 11 Petrography, microscopic fluorescence characteristics and homogenization temperatures of hydrocarbon and aqueous inclusions in the Cretaceous reservoir of well BZ102

(a, b)—第 I 类油包裹体组合发育于石英愈合裂隙, 6762.8 m, (a): 紫外光 (UV) 下呈蓝绿色荧光, (b): 单偏光镜 (-) 下无色; (c, d)—第 I 类油包裹体组合发育于石英次生加大边与碎屑颗粒之间及石英颗粒愈合裂隙中, 6771.6 m, (c): 紫外光 (UV) 下呈蓝绿色荧光, (d): 阴极发光 (CL); (e, f)—第 II 类烃包裹体发育于石英愈合裂隙中且切穿石英次生加大边, 6761.0 m, (e): 紫外光下 (UV) 呈蓝色荧光, (f): 阴极发光 (CL); (g, h)—第 II 类烃包裹体发育在穿石英颗粒裂纹, 6775.5 m, (g): 紫外光 (UV) 下呈淡黄色、蓝色荧光, (h): 单偏光镜 (-) 下无色; (i)—第 I 类烃包裹体显微荧光光谱特征; (j)—第 II 类烃包裹体显微荧光光谱特征; (k)—流体包裹体均一温度 (T_h) 分布直方图

(a, b)—photomicrographs of type I oil inclusion assemblages occurring in annealed microfractures within quartz grains, 6762.8 m, (a) under UV light showing oil inclusions with blue-greenish fluorescence colours, (b) under plain light; (c, d)—photomicrographs of type I oil inclusion assemblages occurring in overgrowth bands and annealed microfractures in quartz grains, 6771.6 m, (c) under UV light showing oil inclusions with blue-greenish fluorescence colours, (d) under cathodoluminescence light (CL); (e, f)—photomicrographs of type II oil inclusion assemblages occurring in annealed microfractures that cut through overgrowth bands of quartz, 6761.0 m, (e) under UV light showing oil inclusions with blue fluorescence colour, (f) under cathodoluminescence light (CL); (g, h)—photomicrographs of type II oil inclusion assemblages occurring in annealed microfractures that cut through quartz grains, 6775.5 m, (g) under UV light showing oil inclusions with pale yellowish and blue fluorescence colours, (h) under plain light; (i)—fluorescence spectral characteristics of type I oil inclusions; (j)—fluorescence spectral characteristics of type II oil inclusions; (k)—homogenization temperature (T_h) histogram of hydrocarbon and coeval aqueous inclusions

裹体的均一温度(T_h)分布在 $45\sim 65^\circ\text{C}$ 之间,与其伴生的盐水包裹体均一温度主要分布在 $95\sim 105^\circ\text{C}$ 之间;第II类油包裹体均一温度分布范围较广,在 $25\sim 85^\circ\text{C}$ 之间,存在明显的超压捕获特征,与其伴生的盐水包裹体 T_h 值在 $115\sim 145^\circ\text{C}$ 之间(图11k)。

2.3.2 油包裹体捕获时的古温度与古压力

在对两种类型油包裹体均一温度、气液比及其伴生盐水包裹体均一温度精确测定的基础上,

根据相平衡原理,利用压力-体积-温度(PVT)模拟软件,可建立流体包裹体的 P - T 相图和等容线,油包裹体及其伴生盐水包裹体等容线的交点即为该类油包裹体捕获时的古温度与古压力。据此,模拟了博孜102井两种类型油包裹体捕获时的古温压特征,结果表明,第I类油包裹体捕获时的古温度为 122°C ,古压力为 56 MPa (图12a);第II类油包裹体捕获时的古温度为 135°C ,古压力为 68 MPa (图12b)。

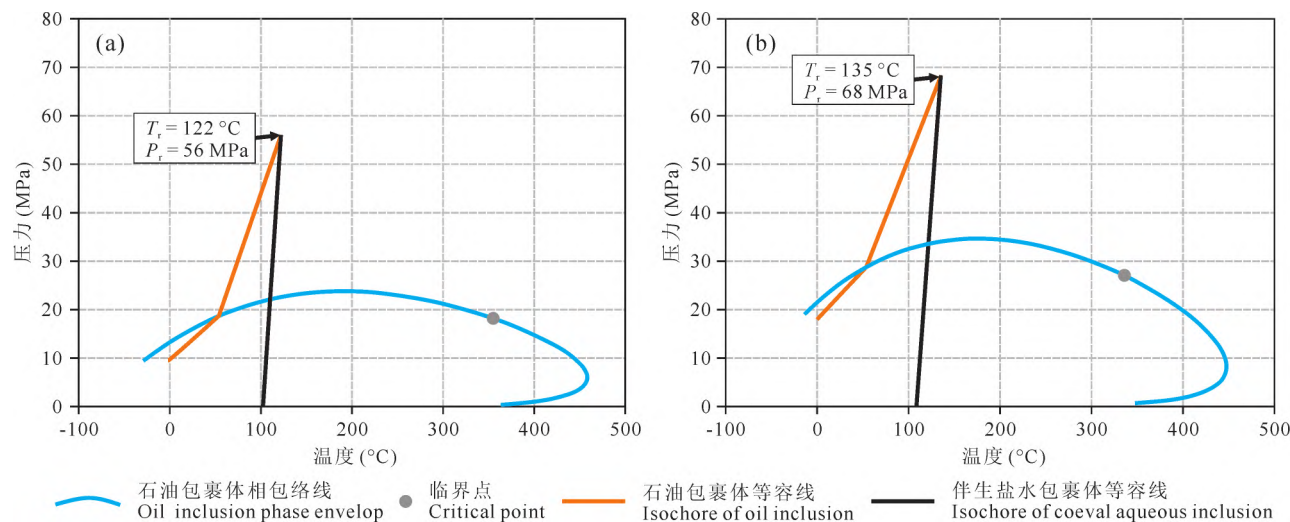


图12 博孜102井两类油包裹体PVT模拟结果

Fig. 12 PVT simulation results of two type oil inclusions in well BZ102

(a) — 第I类油包裹体捕获温度、捕获压力的 P - T 相图; (b) — 第II类油包裹体捕获温度、捕获压力的 P - T 相图

(a) — P - T phase diagram of the type I oil inclusion; (b) — P - T phase diagram of type II oil inclusion, both showing the trapping temperature and pressure

2.3.3 油气成藏模式

单井埋藏史-热史模拟结果表明,博孜地区具有“长期浅埋、晚期快速深埋”的构造-埋藏演化特征(图13),白垩系自沉积以来,经历了长期浅埋藏阶段,地层温度持续处在 80°C 以下;新近纪以来,尤其是康村组沉积之后,博孜地区地层快速埋藏,温度快速升高(图13a)。利用PVT模拟得到的两种类型油包裹体捕获时的古温度和古压力数据,结合单井埋藏史-热史开展油气充注时期厘定,得到的结果更接近实际地质情况,分析结果表明,第I类油包裹体在约 5 Ma 时期被捕获,即早期原油在约 5 Ma 时期充注进来,此时目的层埋深($\sim 4500\text{ m}$)已接近深层;第II类油包裹体的形成时间约为 3 Ma ,指示晚期天然气大规模充注发生在约 3 Ma 以来,此时储集层埋深在约 5800 m 的深层(图13b)。

综合分析认为,博孜地区白垩系现今超深层天然气藏的形成主要发生在深层至超深层,“长期浅埋、晚期快速深埋”的构造-埋藏演化背景是该类型深层油气藏形成的关键。

3 结论

含油气盆地深层—超深层油气资源丰富,经历了盆地演化全过程,存在多类型烃源,但生烃潜力有限,储集层普遍致密,油气运移难度大。从物质与能量守恒、盆地构造-沉积演化与生排烃史、储层地质流体分析等多个角度,探讨了形成现今深层—超深层油气藏的两种较为可能的成藏形式(模式):中—浅层成藏、深埋保持型;长期浅埋、晚期快速深埋成藏型。以塔里木盆地台盆区碳酸盐岩和库车坳陷碎屑岩为例,证实了两种深层油气成藏模式的合理性。塔里木盆地顺北地区奥陶系深层油气藏为“早期

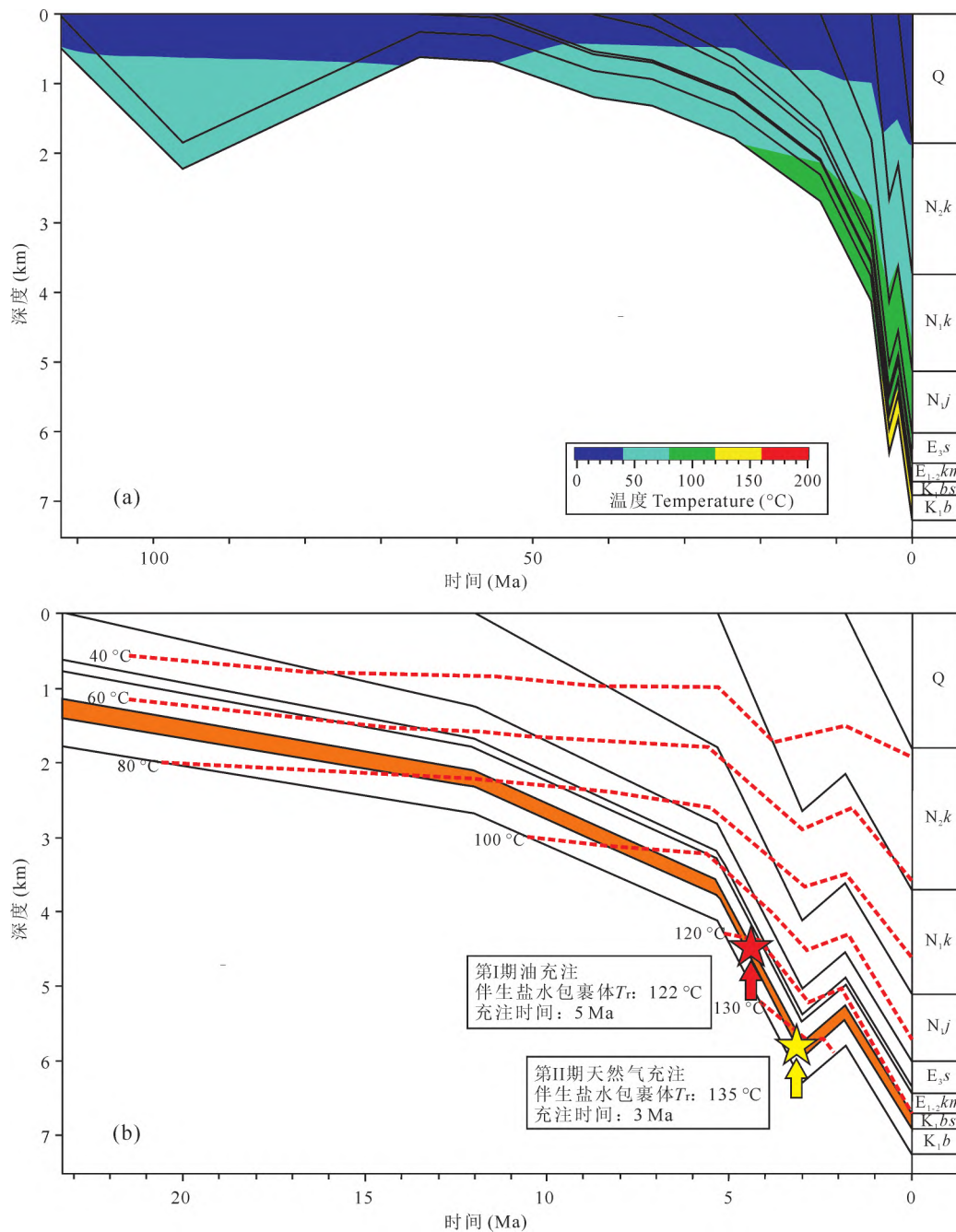


图 13 博孜 102 井埋藏史-热史模拟结果(a)及油气充注时期厘定(b)

Fig. 13 Burial-thermal history (a) and hydrocarbon charge timing (b) of well BZ102 in the Kuqa depression

中—浅层油成藏、后期深埋持续保存”的成藏模式，油气成藏后相对稳定的构造背景是油气藏能保持至今的关键因素；库车坳陷博孜地区白垩系现今超深层凝析气藏的形成主要发生在深层至超深层，“长期浅埋、晚期快速深埋”的构造-埋藏演化背景是该类型深层油气藏形成的关键。

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Deep petroleum accumulation models revisited: Case studies from the Tarim basin

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Abstract

The progressive expansion of oil and gas exploration to the deep and ultra-deep strata suggests a great potential in deep parts of petroliferous basins for oil and gas exploration. A number of scientific issues have also emerged including hydrocarbon generation, reservoir development and hydrocarbon accumulation in the deep and ultra-deep strata. On the basis of extensive literature review, this paper elaborates upon the unique environments for deep oil and gas accumulations and factors controlling the deep source, reservoir and cap rock formation, and focuses on four fundamental issues that need to be paid more attention in studying the formation and evolution of deep and ultra-deep oil and gas reservoirs: ① research on oil and gas generation, reservoir storage space formation, oil and gas phase behavior and hydrocarbon migration in deep strata needs to be carried out on the premise of the basic laws of material balance and energy conservation; ② the deep strata has experienced the entire basin evolution process from shallow to deep; it is thus necessary to study oil and gas accumulation from the perspective of dynamic evolution; ③ it is necessary to explore new experimental methods and enhance our understanding on the stability of liquid hydrocarbons in deep basin and their maximum preservation depth (temperature); ④ more attention needs to be paid to the integration of multidiscipline and intersecting multiple technologies to unravel deep and complex geological issues. Two possible reservoir forming models for deep and ultra-deep reservoirs are proposed: ① middle-shallow depth emplacement and deep-burial preservation type; ② prolonged shallow burial and rapid late-stage deep burial (condensate) gas accumulation type. Taking the Shunbei area in the central carbonate platform of the Tarim basin and the Bozi-Dabei area in the Kuqa depression as two typical examples, oil and gas accumulation in the two areas were studied using reservoir geochemical analysis, a suite of fluid inclusion analysis techniques, *in-situ* calcite U-Pb geochronology and basin modelling. The results show that the Ordovician deep oil and gas reservoirs in the Shunbei area are characterized by “an early middle-shallow accumulation, a late-stage deep burial and sustained preservation”. Relatively stable structural settings after oil and gas being entrapped is the key factor for oil and gas reservoirs to be preserved. The formation of the Cretaceous ultra-deep condensate gas reservoirs in the Bozi area mainly occurred in deep- to ultra-deep settings. A “prolonged shallow burial and late-stage rapid deep burial” is crucial to the formation of this type of deep condensate gas reservoirs. Both examples attest the two reservoir forming models proposed for deep oil and gas reservoirs. The findings may have significant implications for enhancing our understanding of deep and ultra-deep oil and gas accumulations and may provide new insights for deep and ultra-deep oil and gas exploration elsewhere.

Key words: deep petroleum accumulation models; hydrocarbon charge geochronology; fluid inclusions; basin modeling; Tarim basin